

若尔盖高原不同深度泥炭在增温条件下CH₄释放*

刘亮锋¹ 陈槐^{2**} 彭长辉¹ 朱求安¹ 李宝鑫³

¹西北农林科技大学黄土高原土壤侵蚀与旱地农业国家重点实验室 杨凌 712100

²中国科学院成都生物研究所山地生态恢复与生物资源利用重点实验室 成都 610041

³青海师范大学科学技术处 西宁 810008

摘要 不同深度泥炭所处的环境不同,对甲烷释放的影响也不同,温度是影响土壤甲烷释放的重要环境因子之一。为了解温度变化对不同深度特别是深层泥炭甲烷释放的影响,通过模拟实验监测不同深度泥炭甲烷释放对增温(8和18℃)的响应特征,并分析土壤可溶性有机碳(DOC)和微生物量碳(MBC)对甲烷释放的影响。结果显示,温度升高导致泥炭地总甲烷排放量降低,8℃时甲烷的总释放量(以CH₄-C计)为(4.94 ± 0.70) mg m⁻² d⁻¹,18℃时为(3.16 ± 0.69) mg m⁻² d⁻¹。不同深度泥炭甲烷释放对温度升高的响应不同,表层泥炭甲烷释放随温度的增加而增加,深层泥炭甲烷释放随温度增加而降低,这可能受不同深度土壤有机质所含碳种类的影响。低温时DOC抑制甲烷释放,增温时DOC对甲烷释放没有影响,这可能是受DOC来源的控制。甲烷释放随MBC含量的增加而降低。综上所述随着气候变化,温度升高,若尔盖高原泥炭地甲烷的释放量可能降低。(图6表1参65)

关键词 泥炭地;甲烷;培养试验;增温;可溶性有机碳;微生物量碳

CLC X171 (27)

CH₄ emissions under warming schemes from peatlands of different depths in the Zoige Plateau*

LIU Liangfeng¹, CHEN Huai^{2**}, PENG Changhui¹, ZHU Qiu'an¹ & LI Baoxin³

¹State Key Laboratory of Soil Erosion and Dryland Farming on the Loess Plateau, College of Forestry, Northwest A&F University, Yangling 712100, China

²Key Laboratory of Mountain Ecological Restoration and Bioresource Utilization & Ecological Restoration Biodiversity Conservation Key Laboratory of Sichuan Province, Chengdu Institute of Biology, Chinese Academy of Sciences, Chengdu 610041, China

³Department of Science and Technology, Qinghai Normal University, Xining 810008, China

Abstract With the warming of climate, much degradation has taken place in peatlands on Zoige Plateau. But its effect on the CH₄ emission at soil of different depths remains unknown. In this study, soil of different depths from the Zoige Plateau was incubated at two temperatures (8 °C and 18 °C) to detect CH₄ emission; the effects of dissolved organic carbon (DOC) and microbe biomass carbon (MBC) on CH₄ emission were also determined. The results showed that enhanced temperature decreased CH₄ emission from 4.94 ± 0.70 mg m⁻² d⁻¹ at 8 °C to 3.16 ± 0.69 mg m⁻² d⁻¹ at 18 °C. The peat at different depths showed different responses to warming. Probably due to the varying soil substrate quality among the whole profile, CH₄ emission was increased at the surface but decreased at deeper layers. We also found that DOC had negative effect on CH₄ emission at 8 °C but no effect at 18 °C, which should be correlated to the origin of DOC. MBC had a negative effect on CH₄ emission. Based on the results, we can conclude that CH₄ emission may decrease with the warming of Zoige Plateau peatlands.

Keywords peatlands; methane; incubation experiment; warming; DOC; MBC

全球主要泥炭地自全新世以来积累了大约600 Pg C, 储存的有机碳占陆地总有机碳的1/3。作为陆地最大的有机碳库^[1-2], 泥炭地在全球碳循环中占有重要的地位, 在预测区域

甲烷释放过程中也具有重要的作用^[2-3]。甲烷是一种温室效应比CO₂高15-30倍的温度气体^[4-5]。研究发现全球泥炭地向大气释放甲烷的速率为0.03 Pg a⁻¹^[3, 6]。以升温为主的气候变化对全球各种生态系统产生了很大的影响, 特别是对气候环境特殊的泥炭地^[7-9]。温度升高加速了泥炭地有机碳的分解, 也增加了泥炭地向大气释放温室气体的速率^[10-11]。所以研究泥炭地甲烷释放与全球气候变化之间的关系对预测未来气候变化具有重要的意义。

泥炭地水位、植被类型、降水和温度等对甲烷释放都有很大的影响^[12-15]。其中温度对泥炭地甲烷释放的影响最为复

收稿日期 Received: 2015-05-15 接受日期 Accepted: 2015-06-15

*中国科学院“百人计划”项目、四川省“千人计划”项目、中国科学院对外合作项目(151751KYBS20130027)和青海师范大学基础研究基地开放课题资助 Supported by the 100 Talents Program of the Chinese Academy of Sciences, the 1000 Talents Program of Sichuan Province, the External Cooperation Program of Chinese Academy of Sciences (151751KYBS20130027) and the Open Project from Fundamental Research Base of Qinghai Normal University.

**通讯作者 Corresponding author (E-mail: chenhuai@cib.ac.cn)

杂^[16]。温度可以通过影响土壤物理化学性质影响甲烷释放,研究表明温度变化影响土壤湿度,并对甲烷释放产生促进或者抑制的作用^[17-18];温度也通过影响土壤中微生物群落结构对甲烷释放产生影响^[19]。有些学者发现升温促进了泥炭地甲烷的释放^[20-22],也有研究发现升温对泥炭地甲烷释放具有抑制作用^[23]。泥炭地不同深度泥炭是不同历史时期气候环境下的植被残体或分泌物,不同历史时期气候环境有很大的差异,使泥炭地不同深度的泥炭类型、微生物群落结构及理化性质有很大的差异^[24]。所以在温度升高时不同深度泥炭甲烷释放可能有很大的差异。

若尔盖高原形成的泥炭地面积已有4 605 km²,积累的碳接近0.48 Pg,泥炭平均深度达到1.39 m^[25]。高水位形成的厌氧和低温环境是泥炭积累的主要原因^[1, 26-27]。泥炭地对环境变化相对于其他生态系统更为敏感,特别是处于厌氧低温保护状态下的深层泥炭^[10]。研究表明过去40年该地区的平均温度上升了0.2-0.4^[28]。虽然泥炭地甲烷释放对增温的响应已经有很多的研究^[28-31],但深层泥炭甲烷释放对温度变化的研究还很少。因此,本研究通过培养实验测定不同深度泥炭甲烷释放对增温的响应,同时分析土壤可溶性有机碳(DOC)和微生物量碳(MBC)在培养过程中对甲烷释放的影响,以

期为若尔盖高原在气候变化过程中的区域气候预测提供重要的数据支持。

1 材料与方法

1.1 研究区概况

研究样地位于四川省红原县瓦切乡(33°06'25" N, 102°38'33" E),属于若尔盖高原典型泥炭地(图1),是我国重要的高寒湿地。年平均气温1.7,最低气温可达-10,最高大约为9.1-11.4。有生长季和非生长季之分,生长季短(6-10月),常年处于漫长的非生长季(10月至次年6月)。年降水量为640 mm,大多集中在生长季。若尔盖高原除了泥炭地,还有面积更广的亚高山草甸等。植被主要以藏嵩草(*Kobresia tibetica*)和木里苔草(*Carex muliensis*)为主。

1.2 样品采集

样品采集于2013年10月份(生长季末期)。采样前用刻刀除去地表的植被。用直径10 cm土钻随机取3个100 cm泥炭柱,每个泥炭柱以10 cm为一个深度分割,然后用铝箔纸包裹带回实验室,在输送过程中尽量保证泥炭原有土壤结构的完整性。在实验室,每个深度的泥炭分为两部分:一部分除去其

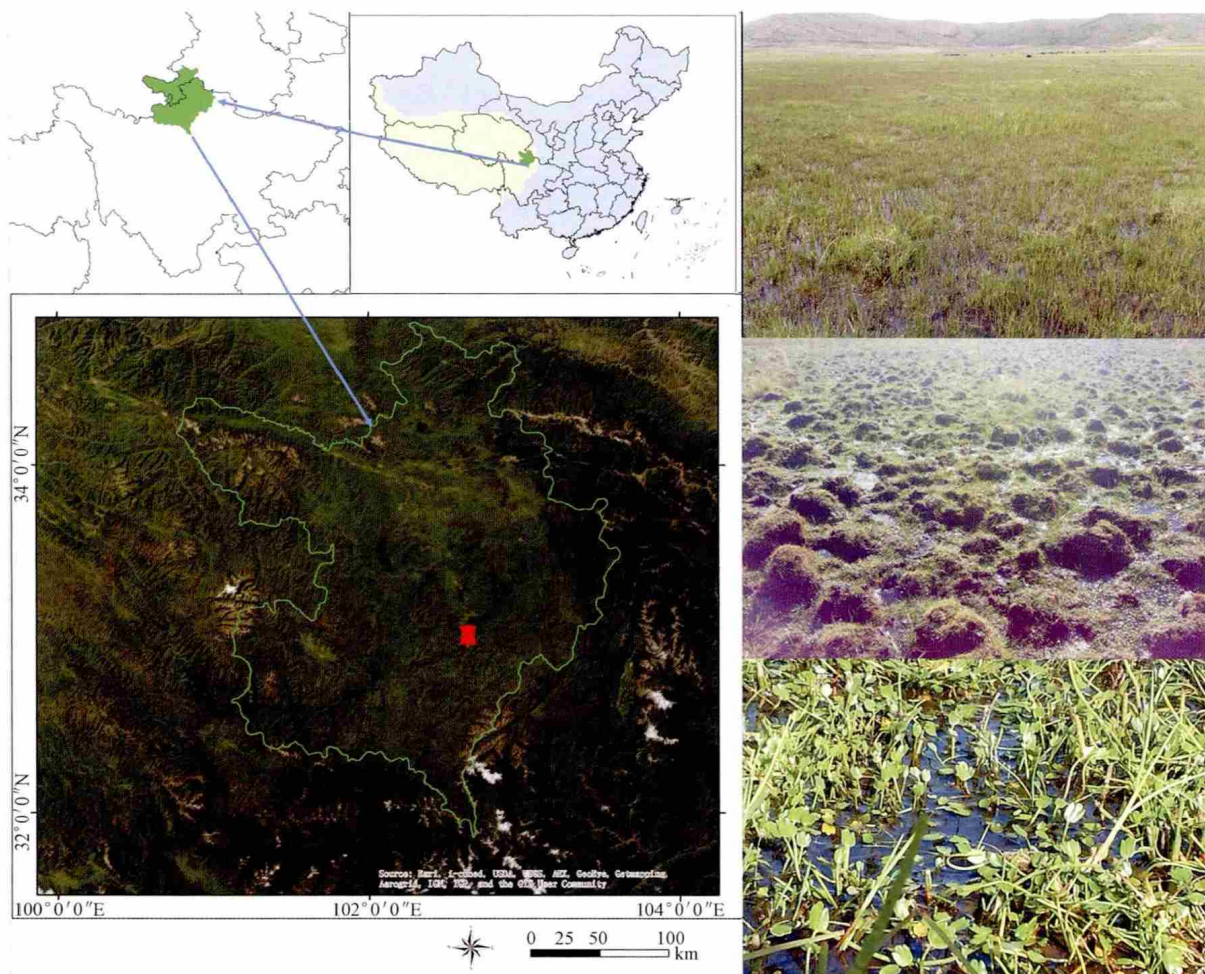


图1 采样点位置示意图。

Fig. 1 The site of sampling.

中可见的植物根系(< 1 mm)及石子,用于培养试验;一部分过2 mm筛子,用于物理化学性质测定。

1.3 培养实验

不同深度土壤(0-100 cm,每10 cm为一个深度)分别培养在18 和8 下。取60 g泥炭放置在500 mL培养瓶中,顶空气体用N₂置换为厌氧状态,用聚乙烯瓶塞保持瓶子密封,然后放置在相应的温度下培养。培养过程中土壤含水量保释恒定。培养开始后每过24 h用带有三通阀的注射器取10 mL顶空气体,用气相色谱(Agilent 7890A, Agilent Co., USA)测定其中甲烷浓度。每次取样之后用N₂再次置换顶空气体,除去瓶中积累的甲烷后继续培养,共培养35 d。甲烷释放速率的计算公式为:

$$F = \frac{PPM \cdot M_0 \cdot T_0 \cdot P}{2.4 \cdot T \cdot P_0} \cdot \frac{V_0}{A} \cdot \frac{1}{d}$$

其中 F (mg m⁻² d⁻¹)为CH₄的释放速率; M_0 为CH₄的相对分子质量; T 和 P 分别为培养瓶中顶空气体的温度和大气压; T_0 和 P_0 分别为标准状态下的温度和大气压; V_0 为瓶子顶空气体的体积,为500 × 10⁻⁶ m³; A 为通过土壤容重计算的土壤释放CH₄有效面积(m²); d 为土壤培养天数,为1 d。

1.4 土壤理化性质分析

总碳测定:将风干土壤过200目筛,然后用总碳分析仪(LIQUIL TOCH, Elementar, Germany)测定土壤总碳含量。

总氮测定:凯氏定氮法。首先将风干并过60目筛的土壤在硫酸钾-硫酸铜为催化剂的浓硫酸环境中,高温煮沸3 h,然后用凯氏定氮仪(Kjeltec™8400 Analyzer Unit Unit Unit, Foss, Sweden)测定总氮含量。土壤碳氮比为总碳和总氮含量的比值。

DOC浓度测定:将未风干土样过2 mm筛,然后用碳酸氢钠和水1:10摇床震荡下提取,用0.45 μm的滤膜过滤,得到滤液用连续流动分析仪(SKALAR San++, SKALAR Co., Netherlands)测定其中DOC浓度。

MBC浓度测定:氯仿熏蒸法^[32-33]。

DOC和MBC浓度在培养前后都进行测定,以计算培养过程中DOC和MBC的变化。

1.5 数据分析

用SPSS 20.0统计分析软件和Microsoft Office Excel 2007对数据进行分析。具体分析方法是通过对方差分析比较甲烷释放随深度的变化。通过相关性分析比较DOC浓度和MBC浓度对甲烷释放量的影响。分析之前对所有数据进行正态分布检验。

2 结果

2.1 土壤物理化学性质

土壤总碳和碳氮比随着深度有显著的差异($P < 0.05$;表1),总碳含量最高为(551.37 ± 24.62) g/kg,最低为(171 ± 21.51) g/kg,碳氮比的变化范围为12.36-35.39。随着深度的增加,DOC浓度没有显著变化($P > 0.05$;表1)。MBC浓度随着深度的增加逐渐降低($P < 0.05$;表1)。

表1 取样点不同深度泥炭理化性质

Table 1 Origin soil properties of different depths in the sampling site

深度 Depth (δ/cm)	总碳 TC (w/g kg ⁻¹)	TC/TN	DOC (w/mg kg ⁻¹)	MBC (w/g kg ⁻¹)
0-10	491.93 ± 4.74b	26.35 ± 0.95b	337.93 ± 8.02a	2.67 ± 0.10a
10-20	400.4 ± 46.44c	24.44 ± 1.70bc	261.80 ± 26.09b	2.68 ± 0.13a
20-30	211.85 ± 7.4f	14.14 ± 0.36e	197.63 ± 1.76def	2.08 ± 0.08c
30-40	171.49 ± 21.54g	12.36 ± 1.00e	162.03 ± 5.5f	1.70 ± 0.04d
40-50	297.59 ± 5.82e	22.48 ± 0.35c	178.83 ± 4.15ef	2.06 ± 0.17c
50-60	259.68 ± 8.85e	17.06 ± 0.62d	174.90 ± 6.82f	1.74 ± 0.06d
60-70	476.63 ± 1.35b	33.43 ± 0.25a	231.63 ± 8.38bcd	2.35 ± 0.01b
70-80	347.73 ± 18.60d	27.81 ± 1.39b	242.9 ± 7.64bc	1.68 ± 0.05d
80-90	551.37 ± 24.62a	35.39 ± 0.72a	236.97 ± 2.37bc	1.55 ± 0.03d
90-100	268.83 ± 31.61e	21.29 ± 1.20c	215.6 ± 5.78cde	2.54 ± 0.01ab

数据为平均值 ± 标准差。小写字母表示不同深度中相应变量的差异性;TC/TN:总碳与总氮比;DOC:可溶性有机碳;MBC:微生物量碳。

Values are means ± standard error. Different letters among depths indicate significant differences at $P < 0.05$. TC: total carbon; TN: total nitrogen; DOC: dissolved organic carbon; MBC: microbe organic carbon.

2.2 甲烷释放异质性

2.2.1 增温对甲烷释放的影响 两个温度下甲烷释放有显著的差异($P < 0.05$)。在8 时甲烷的释放总量(以CH₄-C计)为(4.94 ± 0.70) mg m⁻² d⁻¹,在18 时甲烷释放总量为(3.16 ± 0.69) mg m⁻² d⁻¹。相对于8 ,18 时甲烷释放降低了36%(图2)。

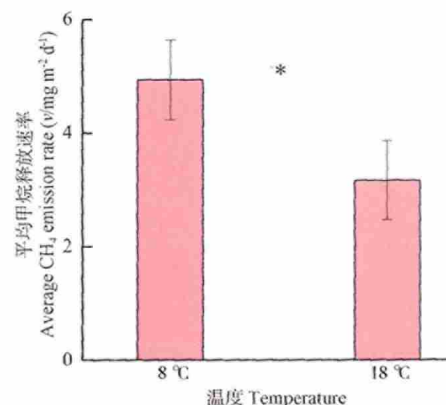


图2 不同温度下甲烷释放($P < 0.05$)。

Fig. 2 CH₄ emission under different temperatures.

2.2.2 甲烷释放随深度的变化 甲烷释放随深度有显著变化($P < 0.05$),但两个温度梯度下甲烷随深度的变化不同。在8 时,甲烷释放随深度的差异极显著($P < 0.01$;图3A),其中表层0-20 cm甲烷的释放量显著低于其他深度甲烷释放量,随着深度增加甲烷释放增加,70-90 cm处甲烷的释放量高于其他深度,90-100 cm处又降低。在18 时,随着深度变化甲烷释放没有显著的差异($P > 0.05$;图3B),其中0-20 cm处甲烷释放量高于其他深度甲烷的释放,20-30 cm处甲烷的释放最低。

2.3 DOC和MBC含量与甲烷释放之间的关系

DOC浓度在培养过程中升高(图4A)。相比8 ,DOC浓度在18 时的增量更大。DOC浓度的变化随着深度有显著差异($P < 0.05$)。MBC含量的变化在两种温度下不同(图4B),在8 时,MBC浓度基本保持不变,部分深度甚至降低,在18 时,MBC浓度比原土有显著增加。MBC浓度变化随着深度没有显著差异($P > 0.05$)。

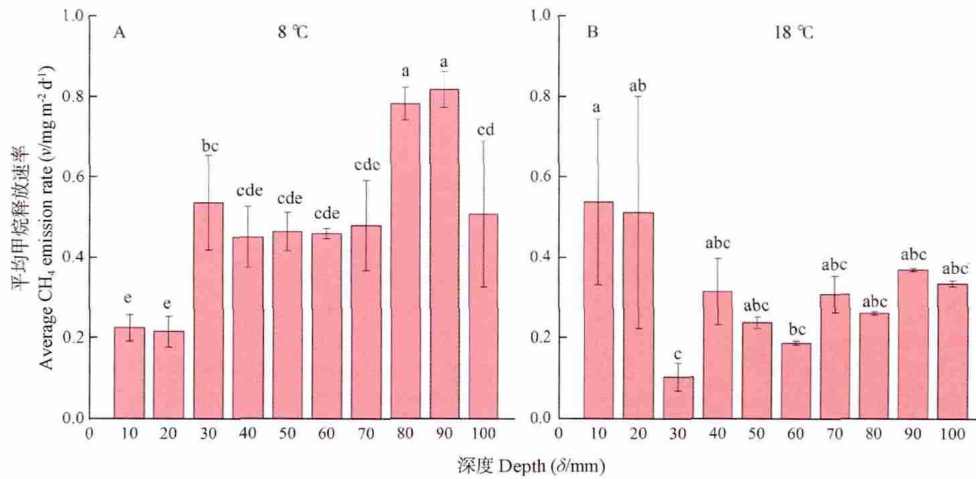


图3 不同深度泥炭在8 (A)和18 (B)培养时的甲烷释放。
Fig. 3 CH₄ emission from different depths at 8 °C (A) and 18 °C (B).

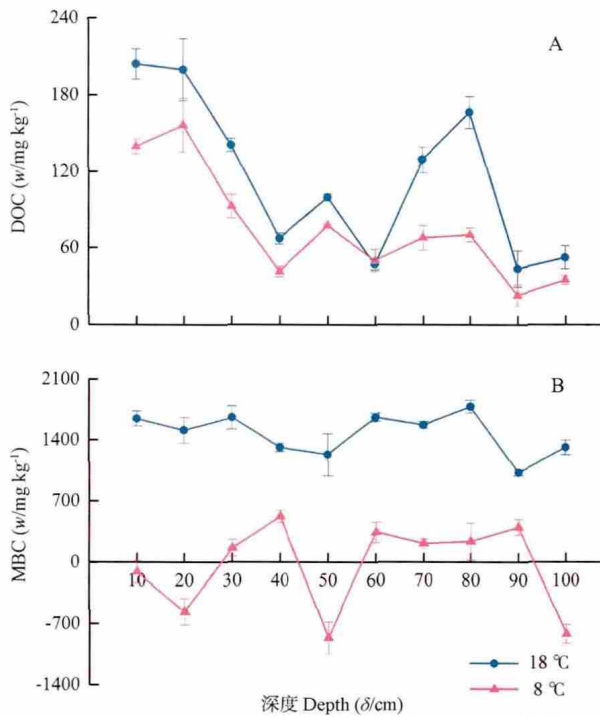


图4 培养过程中DOC浓度(A)和MBC浓度(B)的变化。
Fig. 4 DOC (A) and MBC (B) concentration variation at two temperatures during incubation.

通过分析发现,不同温度下DOC浓度大小对甲烷释放的影响不同。在8 °C时,甲烷释放与DOC浓度之间有很强的负相关($R^2 = 0.45, P < 0.01$;图5A),但是在18 °C时,甲烷释放和DOC的浓度之间没有相关性($P > 0.05$;图5B)。在两种温度下,MBC浓度和甲烷释放之间呈负相关($R^2 = 0.22, P < 0.01$;图6)。

3 讨论

3.1 温度对甲烷释放的影响

本研究发现温度升高使土壤甲烷释放降低,这与很多

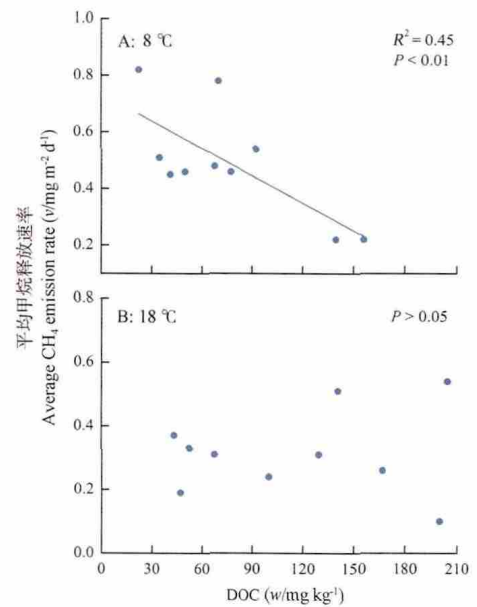


图5 甲烷释放和DOC浓度之间的关系。
Fig. 5 The relationship between CH₄ emission and DOC concentration.

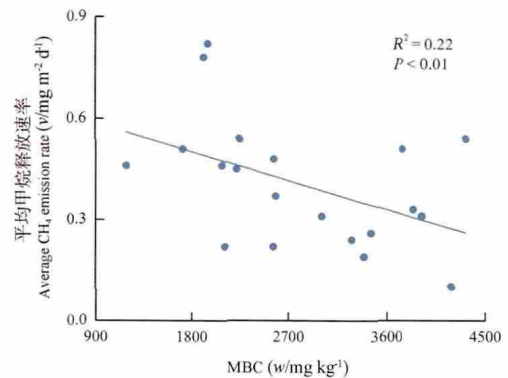


图6 甲烷释放和MBC浓度之间的关系。
Fig. 6 The relationship between CH₄ emission and MBC concentration.

实验研究结果相反。Yang等(2014)在模拟泥炭地退化过程中发现随着温度升高,土壤甲烷释放增加了28%,部分研究发现温度增加时土壤甲烷释放增加了75%-80%^[11, 22]。产甲烷菌的活性会随着温度升高而增强,丰度也增加,促进了土壤甲烷的释放。Eriksson等(2010)发现在温度增加时土壤甲烷释放降低30%^[34]。土壤有机质种类对CH₄释放有直接的影响^[35]。Nilsson等(2009)研究表明土壤有机碳的分解程度对有机碳分解为CH₄或者CO₂具有影响^[36-37]。有机碳的种类影响其分解的方向^[34]。土壤中难分解的高分子化合物分解产生的小分子化合物更多地被分解为CO₂,减少土壤甲烷的释放^[37],易矿化的有机碳分解的小分子化合物易被分解为甲烷^[38-40]。在若尔盖高原泥炭地,深层的碳大多为难分解有机碳^[10],产生的小分子化合物不易被产甲烷菌利用,间接抑制甲烷的释放。同时泥炭地退化过程中,温度升高也使甲烷氧化菌活性增加^[41],导致甲烷释放降低。

3.2 不同深度甲烷的释放

研究发现甲烷释放随深度有显著差异,泥炭基质的物理化学性质随深度的显著变化是甲烷释放随深度变化的主要原因^[42]。总体而言,表层土壤中含有当年植被凋落物及根系分泌物,使得表层土壤有机碳比深层的易分解^[43],会被分解产生更多的甲烷^[38-40]。由于甲烷的释放是甲烷产生和氧化之间的平衡^[34],表层土壤通过植物及土壤空隙与大气进行气体交换^[44-46],增加了表层土壤中甲烷氧化菌的丰度,表层土壤产生的甲烷更多地被氧化,使最终释放的甲烷量降低。深层土壤产甲烷菌长期在受温度限制的环境中,胁迫环境中微生物大多以休眠体的形式存在^[27],当土壤处于相对适宜的环境中时,土壤中产甲烷菌活性增加,同时深层土壤一直处于厌氧环境中,甲烷氧化菌丰度低,整体使深层土壤在环境适宜时甲烷释放增加。所以在8℃时,表层土壤甲烷释放量显著低于深层。温度增加为18℃时,表层土壤甲烷释放增加,深层甲烷释放降低。研究发现产甲烷菌对温度的敏感性高于甲烷氧化菌^[47]。同时深层土壤基质本身不利于甲烷产生,温度升高时更多的难分解有机碳分解,产生的小分子化合物易被分解为CO₂,使深层甲烷释放降低^[37]。

3.3 环境因子对甲烷释放的影响

3.3.1 DOC对甲烷释放的影响 土壤中DOC的浓度是土壤有机碳分解产生DOC和土壤微生物分解消耗DOC之间的平衡^[48-50]。DOC来源不同,对甲烷释放的影响不同^[47]。在低温环境中,土壤中大部分生物代谢处于抑制状态,甲烷大多来源于难分解的有机碳分解产生的DOC,这种DOC对甲烷释放有抑制作用^[37],所以在8℃时,DOC浓度与甲烷释放之间呈负相关性。18℃时,生物代谢加强,很多高分子化合物被分解^[10, 51-52],产生的DOC更多地被分解以CO₂的形式释放^[53-56],对甲烷释放没有很大的影响。所以18℃时,DOC浓度与甲烷产生之间没有相关性。

3.3.2 MBC对甲烷释放的影响 土壤微生物的活性受到温度的影响^[42, 57-58]。土壤有机碳分解是在一系列微生物的作用下完成的,所以甲烷的释放与土壤微生物碳浓度有很强的相关性^[59]。本研究发现微生物量碳与甲烷释放之间呈负相关性。温度增加,使土壤微生物的活性增加,包括产甲烷菌的活性,此时温度的作用增加了土壤甲烷的释放,但由于土

壤基质的不同使甲烷的释放多变^[34-37]。促进甲烷产生的微生物(包括产甲烷菌)是土壤中很多微生物中的极少一部分,温度上升,更多的其他微生物活性和微生物量增加,包括甲烷氧化菌。甲烷氧化菌氧化作用使甲烷释放量降低^[60]。研究发现沼泽地甲烷释放出现波动,大多是甲烷氧化菌活性变化引起的^[61]。另外其他微生物活性和丰度随着温度增加时,会和产甲烷菌等促进甲烷产生的微生物竞争呼吸过程中的电子受体^[23, 62-63],或者产生一些对产甲烷菌有毒害作用的硝酸盐^[64-65],抑制产甲烷菌的活性。整体随着土壤微生物量的增加,甲烷释放降低,呈负相关。

4 结论

本研究通过培养实验发现增温使泥炭土壤甲烷释放降低,但不同深度土壤甲烷释放对增温的响应不同,表层土壤甲烷释放随温度的增加而增加,深层土壤甲烷释放随温度增加而降低,这可能与不同深度土壤有机质的分解程度相关。DOC的来源影响其是否分解为CH₄,易分解有机碳产生的DOC促进甲烷释放,难分解高分子有机碳产生的DOC抑制甲烷释放。MBC浓度随温度增加而增加,但高丰度的微生物群落会对产甲烷菌产生高的生存压力,从而抑制了甲烷的释放。所以随着气候变暖,若尔盖高原泥炭地甲烷释放量可能降低。

参考文献 [References]

- 1 Amaral JA, Knowles R. Methane metabolism in temperate swamp [J]. *Appl Environ Microbiol*, 1994, **60**: 3945-3951
- 2 Baird AJ, Belyea LR, Comas X, Reeve AS, Slater LD. Partitioning litter mass loss into carbon dioxide and methane in peatland ecosystems [J]. *Geophys Monogr Ser*, 2009, **184**: 131-144
- 3 Bak CaF, Conrad R. Competition for electron donors among nitrate reducers, ferric iron reducers, sulfate reducers, and methanogens in anoxic paddy soil [J]. *Biol Fertil Soils*, 1995, **19**: 65-72
- 4 Biasi C, Rusalimova O, Meyer H, Kaiser C, Wanek W, Barsukov P, Junger H, Richter A. Temperature-dependent shift from labile to recalcitrant carbon sources of arctic heterotrophs [J]. *Rapid Commun Mass Spectrom*, 2005, **19**: 1401-1408
- 5 Bridgman SD, Richardson CJ. Mechanisms controlling soil respiration (CO₂ and CH₄) in southern peatlands [J]. *Soil Biol Biochem*, 1992, **24**: 1089-1099
- 6 Chen H, Yang G, Peng C, Zhang Y, Zhu D, Zhu Q, Hu J, Wang M, Zhan W, Zhu E, Bai Z, Li W, Wu N, Wang Y, Gao Y, Tian J, Kang X, Zhao X, Wu J. The carbon stock of alpine peatlands on the Qinghai-Tibetan Plateau during the Holocene and their future fate [J]. *Quaternary Sci Rev*, 2014, **95**: 151-158
- 7 Chen H, Zhu Q, Peng C, Wu N, Wang Y, Fang X, Gao Y, Zhu D, Yang G, Tian J, Kang X, Piao S, Ouyang H, Xiang W, Luo Z, Jiang H, Song X, Zhang Y, Yu G, Zhao X, Gong P, Yao T, Wu J. The impacts of climate change and human activities on biogeochemical cycles on the Qinghai-Tibetan Plateau [J]. *Global Change Biol*, 2013, **19**: 2940-2955

- 8 Chu H, Fierer N, Lauber CL, Caporaso JG, Knight R, Grogan P. Soil bacterial diversity in the Arctic is not fundamentally different from that found in other biomes [J]. *Environ Microbiol*, 2010, **12**: 2998-3006
- 9 Conant RT, Ryan MG, Ågren GI, Birge HE, Davidson EA, Eliasson PE, Evans SE, Frey SD, Giardina CP, Hopkins FM, Hyvönen R, Kirschbaum MUF, Lavallee JM, Leifeld J, Parton WJ, Megan Steinweg J, Wallenstein MD, Martin Wetterstedt JÅ, Bradford MA. Temperature and soil organic matter decomposition rates - synthesis of current knowledge and a way forward [J]. *Global Change Biol*, 2013, **17**: 3392-3404
- 10 Dalva M, Moore TR. Source and sinks of dissolved organic carbon a forested swamp catchment [J]. *Biogeochemistry*, 1991, **15**: 1-19
- 11 Dijkstra FA, Prior SA, Runion GB, Torbert HA, Tian H, Lu C, Venterea RT. Effects of elevated carbon dioxide and increased temperature on methane and nitrous oxide fluxes: evidence from field experiments [J]. *Frontiers Ecol Environ*, 2012, **10**: 520-527
- 12 Ding W, Cai Z. Methane emission from mires and its influencing factors [J]. *Sci Geogr Sin*, 2002, **22**: 619-425.
- 13 Ding W, Cai Z. Effect of temperature on methane production and oxidation in soils [J]. *Appl Ecol*, 2003, **14**: 604-608
- 14 Dorrepaal E, Toet S, Van Logtestijn RSP, Swart E, Van De Weg MJ, Callaghan TV, Aerts R. Carbon respiration from subsurface peat accelerated by climate warming in the subarctic [J]. *Nature*, 2009, **460**: 616-619
- 15 Dutta K, Schuur EAG, Neff JC. Potential carbon release from permafrost soils of northeastern Siberia [J]. *Global Change Biol*, 2006, **12**: 2336-2351
- 16 Eliasson PE, Mcmurtrie RE, Pepper DA, Stromgren M, Linder S, Ågren GI. The response of heterotrophic CO₂ flux to soil warming [J]. *Global Change Biol*, 2005, **11**: 167-181
- 17 Eriksson T, Quist MG, Nilsson MB. Effects of decadal deposition of nitrogen and sulfur, and increased temperature, on methane emissions from a boreal peatland [J]. *J Geophys Res (2005-2012)*, 2010a, **115**: G4
- 18 Eriksson T, Quist MG, Nilsson MB. Production and oxidation of methane in a boreal mire after a decade of increased temperature and nitrogen and sulfur deposition [J]. *Global Change Biol*, 2010b, **16**: 2130-2144
- 19 Feng H, Cheng G, An L. Microbia mediated methane cycle in soils and global change: a review [J]. *J Glaciol Geocryol*, 2004, **26**: 411-419
- 20 Fenner N, Freeman C. Drought-induced carbon loss in peatlands. *Nat Geosci*, 2011, **4**: 895-910
- 21 Fierer N, Craine JM, Mclaughlan K, Schimel JP. Litter quality and the temperature sensitivity of decomposition [J]. *Ecology*, 2005, **86**: 320-326
- 22 Freeman C, Evans C, Monteith DT. Export of organic carbon from peat soils [J]. *Nature*, 2001, **421**: 785-786
- 23 Frolking S, Talbot J, Jones MC, Treat CC, Kauffman JB, Tuittila E-S, Roulet N. Peatlands in the Earth's 21st century climate system [J]. *Environ Rev*, 2011, **19**: 371-396
- 24 Gao J, Ouyang H, Lei G, Xu X, Zhang M, Effects of temperature, soil moisture, soil type and their interactions on soil carbon mineralization in Zoigé alpine wetland, Qinghai-Tibet Plateau [J]. *Chin Geogr Sci*, 2011, **21**: 27-35
- 25 Gauci V, Dise N, Fowler D. Controls on suppression of methane flux from a peat bog subjected to simulated acid rain sulfate deposition [J]. *Global Biogeochem Cycles*, 2002, **16** (4): 1-12
- 26 Gorham E. Northern peatlands: role in the carbon cycle and probable responses to climatic warming [J]. *Ecol Appl*, 1991, **1**: 182-195
- 27 He Z. The methods of soil microbial quantity detection: status and prospects. *Progr Soil Sci*, 1994, **22**: 36-43
- 28 Hicks Pries CE, Schuur EaG. Thawing permafrost increases old soil and autotrophic respiration in tundra: partitioning ecosystem respiration using $\delta^{13}\text{C}$ and $\Delta^{14}\text{C}$ [J]. *Global Change Biol*, 2013, **19**: 649-661
- 29 Huo L, Chen Z, Zou Y, Lu X, Guo J, Tang X. Effect of Zoige alpine wetland degradation on the density and fractions of soil organic carbon [J]. *Ecol Eng*, 2013, **51**: 287-295
- 30 Jenkinson DS, Ladd JN. Microbial biomass in soil: measurement and turnover [D]. *Soil Biochem*, 1981: 415-471
- 31 Joabsson A, Christensen TR, Wallen B. Vascular plant controls on methane emissions from northern peatforming wetlands. *Trends Ecol Evol*, 1999, **14**: 385-388
- 32 Jones DL, Dennis PG, Owen AG, Hees PaWV. Organic acid behavior in soils - misconceptions and knowledge gaps [J]. *Plant Soil*, 2003, **248**: 31-41
- 33 Kato T, Y Tang, S Gu, X Cui, M Hirota, M Du, Y Li, X Zhao, T Oikawa. Carbon dioxide exchange between the atmosphere and an alpine meadow ecosystem on the Qinghai-Tibetan Plateau, China [J]. *Agric For Meteorol*, 2004, **124**: 121-134
- 34 Krüger JP, Leifeld J, Alewell C. Degradation changes stable carbon isotope depth profiles in peatlands [J]. *Biogeosci Discuss*, 2014, **11**: 1383-1412
- 35 Laiho R. Decomposition in peatlands: Reconciling seemingly contrasting results on the impacts of lowered water levels [J]. *Soil Biol Biochem*, 2006, **38**: 2011-2024
- 36 Liang W, Yue J, Wu J, Shi H, Huang G, Liang Z. Seasonal variations of soil microbial biomass: respiration rate and CH₄ emission in black earth rice [J]. *Chin J Appl Ecol*, 2003, **14**: 2278-2280
- 37 Lloyd J, Taylor JA. On the temperature dependence of soil respiration [J]. *Funct Ecol*, 1994, **8**: 315-323
- 38 Mer JL, Roger P. Production, oxidation, emission and consumption of methane by soils: a review [J]. *Eur J Soil Biol*, 2001, **37**: 25-50
- 39 Mills RTE, Dewhirst N, Sowerby A, Emmett BA, Jones DL. Interactive effects of depth and temperature on CH₄ and N₂O flux in a shallow podzol [J]. *Soil Biol Biochem*, 2013, **62**: 1-4
- 40 Munir TM, Perkins M, Kaing E, Strack M. Carbon dioxide flux and net primary production of a boreal tree bog: responses to warming and water-table-lowering simulations of climate change [J]. *Biogeosciences*, 2015, **12**: 1091-1111
- 41 Nedwell DB, Watson A. CH₄ production, oxidation and emission in U.K. ombrotrophic peat bog: influence of SO₄²⁻ from acid rain [J]. *Soil Biol Biochem*, 1995, **27**: 893-903
- 42 Nilsson M, Mikkela C, Sundh GG, Svensson BH, Ranney B. Methane emission from Swedish mires: national and regional budgets and

- dependence on mire vegetation [J]. *Geophys Res*, 2001, **106** (20) : 847-820
- 43 Norton JM, Smith JL, Firestone MK. Carbon flow in the rhizosphere of ponderosa pine seedlings [J]. *Soil Biol Biochem*, 1990, **22**: 449-455
- 44 Parrenin F, DV Masson, P Köhler, D Raynaud, D Paillard, J Schwander, C Barbante, A Landais, A. Wegner, J Jouzel. Synchronous change of atmospheric CO₂ and Antarctic temperature during the Last Deglacial Warming [J]. *Science*, 2013, **339**: 1060-1063
- 45 Roulet NT, Ash R, Quinton W, Moore T. Methane flux from drained northern peatlands: effect of a persistent water table lowering on flux [J]. *Glob Biogeochem Cycles*, 1993, **7**: 749-769
- 46 Roy R, Conrad R. Effect of methanogenic precursors (acetate, hydrogen, propionate) on the suppression of methane production by nitrate in anoxic rice field soil [J]. *FEMS Microbiol Ecol*, 1999, **28**: 49-61
- 47 Sistla SA, Moore JC, Simpson RT, Gough L, Shaver GR, Schimel JP. Long-term warming restructures Arctic tundra without changing net soil carbon storage [J]. *Nature*, 2013, **497**: 615-618
- 48 Strack M, Waddington J. Response of peatland carbon dioxide and methane fluxes to a water table drawdown experiment [J]. *Glob Biogeochem Cycles*, 2007, **21**: 1
- 49 Treat CC, Wollheim WM, Varner RK, Grandy AS, Talbot J, Frolking S. Temperature and peat type control CO₂ and CH₄ production in Alaskan permafrost peats. *Glob Change Biol*, 2014, **20**: 2674-2686
- 50 Tucker CL, Bell J, Pendall E, Ogle K. Does declining carbon-use efficiency explain thermal acclimation of soil respiration with warming [J]. *Glob Change Biol*, 2013, **19**: 252-263
- 51 Turetsky MR, Treat CC, Waldrop MP, Waddington JM, Harden JW, McGuire AD. Short-term response of methane fluxes and methanogen activity to water table and soil warming manipulations in an Alaskan peatland [J]. *J Geophys Res (2005-2012)*, 2008, **113**: G3DOI: 10.1029/2007JG000496
- 52 Van Den Berg LJ, Shotbolt L, Ashmore MR. Dissolved organic carbon (DOC) concentrations in UK soils and the influence of soil, vegetation type and seasonality [J]. *Sci Total Environ*, 2012, **427**: 269-276
- 53 Wang J, Song C, Zhang J, Wang L, Zhu X, Shi F. Temperature sensitivity of soil carbon mineralization and nitrous oxide emission in different ecosystems along a mountain wetland-forest ecotone in the continuous permafrost of Northeast China [J]. *Catena*, 2014, **121**: 110-118
- 54 Wang ZP, Vancleemput O, Baert L. Methane emission and trapment in flooded rice soils as affected by soil properties [J]. *Biol Fertil Soils*, 1993, **16**: 163-168
- 55 Ward SE, Ostle NJ, Oakley S, Quirk H, Henrys PA, Bardgett RD. Warming effects on greenhouse gas fluxes in peatlands are modulated by vegetation composition [J]. *Ecol Lett*, 2013, **16**: 1285-1293
- 56 Whiting GJ, Chanton JP. Primary production control of methane emission from wetlands [J]. *Nature*, 1993, **364**: 794-795
- 57 Winden JFV, Reichart G-J, Mcnamara NP, Benthien A, Damste JSS. Temperature-induced increase in methane release from peat bogs: a mesocosm experiment [J]. *PLoS ONE*, 2012, **7** (6): e39614
- 58 Wright EL, Black CR, Turner BL, Sjögersten S. Environmental controls of temporal and spatial variability in CO₂ and CH₄ fluxes in a neotropical peatland [J]. *Glob Change Biol*, 2013, **19**: 3775-3789
- 59 Yagi K, Minami K. Effect of organic matter application on methane emission from some Japanese paddy fields. *Soil Sci Plant Nutr*, 1990, **36**: 599-610
- 60 Yang G, Chen H, Wu N, Tian J, Peng C, Zhu Q, Zhu D, He Y, Zheng Q, Zhang C. Effects of soil warming, rainfall reduction and water table level on CH₄ emissions from the Zoige peatland in China [J]. *Soil Biol Biochem*, 2014, **78**: 83-89
- 61 Yang J, Tang Y, Xu J. Development and applications of gain-switched fiber lasers [J]. *Photonics Res*, 2013, **1**: 52-57
- 62 Yi S, Wang X, Qin Y, Xiang B, Ding Y. Responses of alpine grassland on Qinghai-Tibetan plateau to climate warming and permafrost degradation: a modeling perspective [J]. *Environ Res Lett*, 2014, **9** (7): 074014
- 63 Yu Z, Loisel J, Brosseau DP, Beilman DW, Hunt SJ. Global peatland dynamics since the Last Glacial Maximum [J]. *Geophys Res Lett*, 2010, **37** (13): 5508-5513
- 64 Zak DR, Kling GW. Microbial community composition and function across an arctic tundra landscape [J]. *Ecology*, 2006, **7**: 1659-1670
- 65 Zhu D, Chen H, Zhu QA, Wu Y, Wu N. High carbon dioxide evasion from an alpine peatland lake: the central role of terrestrial dissolved organic carbon input [J]. *Water Air Soil Pollut*, 2011, **223**: 2563-2569